

1 **The potential of small unmanned aircraft systems and structure-from-motion**
2 **for topographic surveys: A test of emerging integrated approaches at Cwm**
3 **Idwal, North Wales**

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13
14 **Abstract**

15 Novel topographic survey methods that integrate both structure-from-motion (SfM)
16 photogrammetry and small unmanned aircraft systems (sUAS) are a rapidly evolving
17 investigative technique. Due to the diverse range of survey configurations available
18 and the infancy of these new methods, further research is required. Here, the
19 accuracy, precision and potential applications of this approach are investigated. A
20 total of 543 images of the Cwm Idwal moraine-mound complex were captured from a
21 light (< 5 kg) semi-autonomous multi-rotor unmanned aircraft system using a
22 consumer-grade 18 MP compact digital camera. The images were used to produce a

23 DSM (digital surface model) of the moraines. The DSM is in good agreement with
24 7761 total station survey points providing a total vertical RMSE value of 0.517 m and
25 vertical RMSE values as low as 0.200 m for less densely vegetated areas of the
26 DSM. High-precision topographic data can be acquired rapidly using this technique
27 with the resulting DSMs and orthorectified aerial imagery at sub-decimetre
28 resolutions. Positional errors on the total station dataset, vegetation and steep terrain
29 are identified as the causes of vertical disagreement. Whilst this aerial survey
30 approach is advocated for use in a range of geomorphological settings, care must be
31 taken to ensure that adequate ground control is applied to give a high degree of
32 accuracy.

33

34 **Highlights**

- 35 ▪ An integrated sUAS and SfM approach is used to generate high-resolution
36 topographic data.
- 37 ▪ SfM data compared with a total station ground survey.
- 38 ▪ Positional errors on the total station dataset, vegetation and steep terrain are
39 identified as causes of vertical difference between the two datasets.
- 40 ▪ The integration of a combined sUAS and SfM approach is discussed.

41

42 **Keywords**

43 Small unmanned aircraft system, Structure from motion, Digital surface model,
44 Digital elevation model, Topographic surveying

45 **1 Introduction**

46 The use of small unmanned aircraft systems (sUAS) and structure-from-motion
47 (SfM) digital photogrammetry presents a new methodological frontier for topographic
48 data acquisition and is of interest to scientists researching in a range of
49 geomorphological environments (Westoby et al., 2012; Carrivick et al., 2013;
50 Hugenholtz et al., 2013; Tarolli, 2014). Traditionally low-level aerial photography has
51 been acquired using a variety of unmanned platforms including small lighter-than-air
52 blimps, kites, and model fixed-wing and single rotor aircraft (e.g. Wester-
53 Ebbinghaus, 1980; Rango et al., 2009; Smith et al., 2009; Hugenholtz et al., 2013).
54 More recently lightweight (< 5 kg), relatively low-cost multi-rotor aerial platforms have
55 been used to capture low-level imagery (Harwin and Lucieer, 2012; Niethammer et
56 al., 2012; Rosnell and Honkavaara, 2012; Mancini et al., 2013; Lucieer et al., 2014).
57 These sUAS can be programmed to fly semi-autonomously at fixed altitudes along
58 flight lines, ensuring optimal image overlap for digital photogrammetry. A key
59 strength of the integrated sUAS–SfM approach is the degree of automation involved.
60 Previously, a high degree of user experience was a prerequisite for both the
61 operation of aerial platforms and the application of photogrammetric methods to
62 extract meaningful topographic data from aerial imagery (Aber et al., 2010). The
63 premise of SfM as a digital photogrammetric technique is that three-dimensional
64 coordinates can be extracted from sufficiently overlapping photography without the
65 need for camera spatial information (Snavely et al., 2008; Westoby et al., 2012). The
66 integration of SfM with sUAS camera platforms offers a rapid and increasingly cost
67 effective option for geomorphologists to produce digital surface models (DSMs), with
68 resolution and data quality proposed to be on-par with, or better than LiDAR
69 (Carrivick et al., 2013; Fonstad et al., 2013). SfM based topographic surveys have

70 recently been used for a variety of geoscientific applications including quantifying
71 rates of landslide displacement (Lucieer et al., 2013), mapping vegetation spectral
72 dynamics (Dandois and Ellis, 2013), producing DEMs (digital elevation models) of
73 agricultural watersheds (Ouédraogo et al., 2014), quantifying coastal erosion rates
74 (James and Robson, 2012), and measuring rates of glacier motion and thinning
75 (Whitehead et al., 2013). The potential of SfM to aid geomorphological mapping,
76 derive measurements of landforms (morphometry) and quantify geomorphological
77 change is evident. Numerous software packages for SfM are now available and
78 include cloud-based processing, which has the additional benefit of not requiring a
79 high-specification consumer computer capable of handling the image processing.

80

81 Whilst a range of recent studies have sought to quantify data quality and associated
82 error of SfM techniques (Harwin and Lucieer, 2012; Turner et al., 2012; Westoby et
83 al., 2012; Dandois and Ellis, 2013; Fonstad et al., 2013; Hugenholtz et al., 2013;
84 Ouédraogo et al., 2014), further research is beneficial due to the diverse nature of
85 the aerial platforms and consumer-grade digital cameras available for the production
86 of topographic data using this methodology. Existing reports on the effectiveness of
87 integrated multi-rotor based sUAS–SfM approaches describe surveys conducted
88 from relatively low altitudes (< 50 m). The aims of this research are to: (1) provide a
89 systematic account of the data acquisition process associated with this new
90 integrated technique; (2) compare vertical spot heights obtained from the sUAS–SfM
91 survey to those obtained from a total station ground survey; (3) highlight important
92 considerations for researchers seeking to use sUAS and SfM approaches to acquire
93 data for topographic investigations; and (4) provide a baseline for the potential

94 spatial resolutions when using a consumer-grade 18 MP compact digital camera at a
95 target flight altitude of 100 m.

96

97 **2 Study area**

98 The test was undertaken at Cwm Idwal, north Wales ($53^{\circ} 6'50.89''\text{N}$; $4^{\circ} 1'38.38''\text{W}$;
99 Fig. 1) (Appendix A), a large cirque that was last occupied by a glacier during the
100 Younger Dryas Stadial (c.12.9-11.7 ka BP; Bendle and Glasser, 2012). The study
101 area is located on the cirque floor and covers an altitudinal range of ~ 370 to ~ 410
102 m (above ordnance datum). The geomorphology of the site is characterised by a
103 moraine-mound complex ('hummocky moraine') located on both the east and west of
104 Llyn [lake] Idwal (Fig. 1c). These moraines have been the subject of numerous
105 investigations (e.g. Darwin, 1842; Escritt, 1971; Gray, 1982; Addison, 1988; Graham
106 and Midgley, 2000; Bendle and Glasser, 2012) due to their importance for
107 understanding the significance of Younger Dryas glaciers in the British Uplands. The
108 majority of the moraines are 8 to 80 m in length, with the exception being a set of
109 discontinuous stream-breached ridges totalling ~450 m in length which are stacked
110 against the western cirque wall. In places the morphology of the moraines are
111 influenced by glacially abraded bedrock. The prominence of some landforms is also
112 disguised by a peat infill. The southern section of the survey area is characterised by
113 a relatively flat lake infill and steep glacially abraded bedrock slopes. Vegetation on
114 the eastern side of the Cwm is typically restricted to short swards of grass, whereas
115 livestock grazing exclosures erected in the 1950s and 1960s on the western side of
116 Llyn Idwal have promoted the growth of vegetation including a thick cover of
117 common heather (*Calluna vulgaris*), western gorse (*Ulex gallii*), and the occasional

118 rowan (*Sorbus aucuparia*) and silver birch (*Betula pendula*) (Rhind and Jones,
119 2003). A large part of the moraine-mound complex and surrounding area were
120 surveyed with a total station by Graham and Midgley (2000). A similar area was
121 surveyed by a sUAS to allow a direct comparison between total station based data
122 acquisition, and the sUAS–SfM method used for this study.

123

124 **3 Methods and materials**

125 *3.1 Image Image and data acquisition*

126 Aerial imagery was acquired using a Canon EOS-M 18 MP camera suspended from
127 a DJI S800 Hexacopter (Fig. 2). A Photohigher AV130 servo driven gimbal
128 maintained the camera angle close to the nadir. The hexacopter was equipped with
129 a Wookong-M GPS assisted flight controller which allowed for semi-autonomous
130 surveys. Survey flight-lines were pre-programmed via the DJI Ground-Station
131 software package. For all surveys the sUAS was set to a target altitude of 100 m
132 above ground level (AGL) and horizontal ground speed of 2.5 ms^{-1} . The target
133 altitude is calculated in the DJI Ground-Station software using elevation data derived
134 from Google Earth. Parallel flights lines were programmed to have an image sidelap
135 of 80%, whilst taking into account the camera sensor size ($22.3 \times 14.9 \text{ mm}$) and
136 focal length (22 mm). The intervalometer function of the Magic Lantern third-party
137 camera firmware was set to acquire imagery every 2 s along parallel flight lines.
138 Actual image acquisition was every $\sim 4 \text{ s}$, resulting in image capture approximately
139 every 10 m along flight lines. Although image capture can be triggered using the DJI
140 flight controller, an intervalometer was used for its improved reliability and potential
141 to capture excess imagery along flight lines. This allowed for blurred or poor quality

142 imagery to be removed whilst ensuring that an image onlap in excess of 80% was
143 maintained. The camera was set to shutter-priority mode and used a 1/1000 s
144 shutter speed. To provide the required image coverage the survey area had to be
145 split between four flights. The sUAS had a flight-time of ~ 14 min whilst carrying its
146 payload (using an 11 Ah, 22.2 V, 6 cell lithium polymer battery). A generous
147 overhead (~ 2 min) was left in order to safely land the sUAS. In the UK unaided
148 visual line of sight (VLOS) has to be maintained whilst operating sUAS (CAA, 2012).
149 Therefore the ground equipment and launch position were moved between flights to
150 allow the sUAS to be easily observed, and manually controlled if necessary.

151

152 The total station dataset was acquired over multiple survey sessions in 1997 and
153 1998 using a Leica TC600 (Graham and Midgley, 2000). An assessment of error for
154 this data set is unavailable. However, measurement accuracies (expressed as
155 standard deviation) for the TC600 are defined by Leica (1997), with distance
156 measurements accurate to $2 \text{ mm} \pm 2 \text{ ppm}$ and angle (horizontal and vertical)
157 measurements to 1.5 mgon . As the original total station dataset was collected for
158 the purpose of characterising the overall shape of the moraine-mound complex,
159 individual points were collected rapidly. Points recorded whilst the prism pole was
160 not perfectly vertical have the potential to result in misregistration between the two
161 datasets. The extent of the resulting error will be exacerbated by slope steepness
162 and the height of the reflector on the detail pole. The SfM dataset was tied into the
163 same arbitrary co-ordinate system and datum through the use of two brass pin
164 benchmarks located on exposed bedrock on the east and west of Llyn Idwal. Point
165 densities for the validation points reach as high as 20 per 100 m² over the moraine-
166 mound complex (Graham and Midgley, 2000). For the sUAS survey, 19 SfM ground-

167 control points (GCPs) were distributed across the survey area (Fig. 3a). White
168 laminated A3 size targets (297 × 420 cm) were used as GCPs and were found to be
169 adequately visible on the aerial imagery. These GCPs were surveyed with a Leica
170 TC407 total station to a precision of ≤ 1 mm and estimated accuracy of < 3 cm.

171

172 *3.2 Image processing and analysis*

173 From the original set of 824 images, 543 images were selected for model
174 reconstruction using the Agisoft Photoscan 1.0.0 (build 1795) software package.
175 Images were visually assessed for quality and blurry images were removed prior to
176 processing. Image processing followed the recommended procedure outlined by
177 Agisoft (2013). Image processing was conducted on a HP Z820 workstation
178 equipped with dual Intel Xeon E5-2690 processors, 128 GB RAM, and nVidia 680
179 graphics card. As GPS information for camera positions were not collected, images
180 were aligned using the 'Generic Pair Preselection' parameter. This parameter
181 detects matching features between images at a lower accuracy first, to reduce
182 overall processing time (Agisoft, 2013). Photoscan provides nominal parameters for
183 setting the target accuracy to which the images are aligned. Here the 'high' setting
184 was used to obtain the best possible image alignment accuracy. Nineteen GCPs
185 were then identified on imagery within the software a total of 674 times, with the XYZ
186 coordinates input for each point. The sparse point cloud was optimised using a
187 marker accuracy of 0.001 m and focal parameters (F_y and F_x) defined in the image
188 headers. Camera radial and tangential distortion coefficients (K_1 , K_2 , K_3 , P_1 and P_2)
189 were automatically estimated by Photoscan. A dense point cloud was then produced
190 using the 'medium' quality setting. Again, this is a nominal setting that relates to the

191 geometric accuracy of the target dense point cloud produced within Photoscan.
192 Aggressive depth filtering was used to remove outliers from the dense point cloud
193 (Agisoft, 2013). The dense point cloud and polygonal mesh was generated using a
194 target point count of 3×10^5 . An additional sparse point cloud and a DSM were
195 produced for comparative purposes. An orthorectified aerial image was produced
196 using the 'orthophoto' and 'mosaic' parameters with colour correction enabled.
197 Where image overlap occurs, the 'mosaic' parameter ensures that images with pixels
198 closest to the image centre are used preferentially for orthophoto generation (Agisoft,
199 2013).

200

201 Data handling and the analysis of geographic data were conducted using
202 QuantumGIS 2.0 and ArcGIS 10.1. SfM height (SfMz) was subtracted from ground
203 height (GSz) for 7761 independently surveyed data validation spot heights derived
204 by total station survey (Fig. 3a) providing a vertical difference. The vertical difference
205 was converted into a raster surface with a 2.1 m cell size using an ordinary kriging
206 function in ArcGIS 10.1. This allowed the vertical difference to be visualised. RMSE
207 (root mean square error) and MD (mean difference) were calculated for the vertical
208 difference (SfMz-GSz). Two zones of contrasting vegetation cover (Z1 and Z2; Fig.
209 3a) were mapped from orthorectified aerial imagery, and used to quantify vertical
210 difference associated with contrasting vegetation types. Z1 is characterised by a
211 continuous ground cover of heather, gorse, and occasional shrub and is located on
212 the western side of Llyn Idwal. Z2 consists of grassland and exposed bedrock, and is
213 also located on the western side of Llyn Idwal.

214

215 **4 Results**

216 The dense point cloud was composed of 31,474,859 unique points. With the
217 exception of the extremities of the model, the effective overlap was > 9 images per
218 point. 'Noisy' anomalies are present where the surface of reflective water-bodies are
219 reconstructed. The orthorectified images had a 0.022 m per pixel resolution, and the
220 DSM as seen in Fig. 3b had a 0.088 m per pixel resolution. These resolutions were
221 achieved from an average flight altitude of 117.282 m AGL as reconstructed from the
222 imagery. Discrepancy between the target flight attitude and actual flight attitude is
223 likely to be caused by the use of low resolution Google Earth elevation data for flight
224 planning, and error associated with the use of barometric pressure sensors for
225 determining relative height (see DJI, 2013). Photoscan reported a total RMSE value
226 of 0.033 m, calculated from the 19 SfM GCPs. The total x and y RMSE values
227 reported by Photoscan were 0.019 and 0.020 m respectively. The total vertical
228 RMSE value was 0.018 m.

229

230 Spot heights ($n = 7761$) from the ground survey (GSz) and DSM (SfMz) are in broad
231 agreement, although the vertical difference is a higher than that reported by the 19
232 SfM GCPs used during the image processing stage. The vertical difference is
233 visualised in Fig. 3c. The dense point cloud provides a vertical RMSE value of 0.517
234 m (Table 1). The differences for the DSM are offset from zero, with a mean
235 difference of 0.454 m. The majority of the height values on the DSM were within the
236 ± 1 m range (99.8%). However, only 55.4% of the SfM DSM values were within \pm
237 0.5 m of the ground survey data. Isolated spot heights were found to be as much as
238 $- 0.705$ m under the actual ground survey (GSz) and as much as 4.347 m over.

239 When vertical RMSE is calculated separately, RMSE for the east (less densely
240 vegetated) is significantly lower (RMSE = 0.200; n = 1988), than the west (RMSE =
241 0.588; n = 5773) with 98.8% of height values for the east falling within the ± 0.5 m
242 range (Fig. 4).

243

244 Two contrasting vegetation zones (Z1 and Z2 in Fig. 3a) were investigated. Z1 had
245 an RMSE value of 0.789 m (n = 244). In contrast, Z2 produced a lower RMSE value
246 of 0.362 m (n = 205). The calculated RMSE values for slopes gentler than 20° and
247 those steeper than or equal to 20° were examined for both patches. The values are
248 0.031 and 0.030 m higher for slopes steeper than 20° regardless of the vegetation
249 type. Where RMSE was calculated for separate 10° bins for the entire dataset (7761
250 observations), excluding the $60\text{--}70^\circ$ bin, the reported RMSE value increases on
251 progressively steeper slopes (0.444 to 0.838; Table 2). The $80\text{--}90^\circ$ bin comprised
252 one observation, which shows a high vertical difference (2.222 m).

253

254 An additional analysis of the DSM derived from the sparse point cloud (2,058,037
255 points) was conducted. The sparse point cloud produced a coarser resolution DSM
256 at 0.258 m per pixel. Unlike the dense point cloud, the sparse point cloud did not
257 produce 'noisy' anomalies related to reflective water-bodies. Points from the SfM
258 DSM and the ground survey data were also in broad agreement with 98.9% of the
259 data within the ± 1 m range, and 58.5% of the data in the ± 0.5 m range. The total
260 vertical RMSE value was 0.505 m. The sparse point cloud derived DSM produced a
261 wider range of outlying values, with minimum and maximum anomalies of -3.416
262 and 3.782 m.

263 **5 Discussion**

264 *5.1 Causes of vertical disagreement*

265 Causes of poor surface representation and vertical disagreement between the two
266 data sets have been investigated and include: (1) vegetation; (2) slope angle; and (3)
267 unintentional random error related to the acquisition of the original total station
268 dataset. Vegetation is a known cause of poor surface representation in DEMs
269 derived from both photogrammetry (Lane, 2000; Marzloff and Poesen, 2009), and
270 airborne LiDAR (Lui, 2008; Spaete et al., 2011; Hladik and Alber, 2012). A visual
271 assessment of high vertical difference against the orthorectified imagery shows that
272 error is particularly pronounced around trees, and in areas vegetated with heather
273 (Fig. 5a). Dense vegetation types obstruct line-of-sight of actual ground level, thus
274 generate a vertical difference between the two datasets (Table 1). This difference
275 generated by vegetation is also apparent when the east (sparsely vegetated) and
276 west (densely vegetated) are visualised together (Fig. 3c) or where RMSE is
277 calculated for the two zones of contrasting vegetation (Z1 and Z2). For the examples
278 of Z1 (heather and other shrubs) and Z2 (grasses and exposed bedrock), the
279 presence of a thick covering of vegetation produces an additional 0.434 m RMSE
280 value (Table 1). Whilst the total station data provides information that can be used to
281 produce a bare earth DEM of the moraines, the data presented from SfM
282 photogrammetry accounts for the surface plus vegetation, and therefore represents a
283 DSM. Fig. 5a exemplifies this error, showing how a ground survey point located
284 under a silver birch generates a vertical difference between the two datasets.
285 Similarly, in other areas of the Cwm Idwal DSM, this problem arises due to tilted
286 bedrock rafts with near vertical and in places overhanging sides (Fig. 5b), generating
287 the outlying vertical difference of 4.347 m. As DSMs are essentially 2.5 dimensional

288 representations of the earth's surface and associated surface features, true 3
289 dimensional representation of overhanging surfaces is not possible (Bernhardsen,
290 2002). If the same SfM approach was applied to un-vegetated terrain (e.g. braided
291 channels in Javernick et al., 2014), a significantly lower degree of vertical difference
292 would be expected.

293

294 Further vertical differences between the two topographic datasets is also likely to be
295 the result of unintentional random errors in the ground survey dataset caused by the
296 reflector detail pole not being held perfectly level during point acquisition. The
297 vertical difference caused by this operational error appears to be exacerbated on
298 steep slopes (Table 2). For example, on a perfectly horizontal surface, if the reflector
299 (with the detail pole set to the minimum high of 1.3 m) was inclined at 10° from
300 vertical opposed to being perfectly vertical, the calculated positional and vertical
301 errors would be 0.226 and 0.020 m respectively. However, if the detail pole was
302 inclined at 10° from vertical on a slope of 30° , the expected vertical error would reach
303 the decimetre range. As 16.3% of the 7761 observations were made on slopes $>$
304 30° , additional errors should be expected. An example where positional
305 misregistration between the two datasets has occurred is presented in Fig. 5c. Here
306 points taken in the vicinity of a steep-sided tilted bedrock raft with near vertical
307 slopes have resulted in vertical disagreement exceeding 1 m. In this circumstance,
308 sub-decimetre positional errors on the ground survey data or poorly resolved
309 features on the SfM DSM promote a high degree of localised vertical disagreement
310 between the two datasets.

311 *5.2 Benefits and practical considerations*

312 The sUAS-SfM technique is in many ways superior to a conventional total station
313 ground survey and performed comparably to a range of recent SfM data validation
314 studies (Table 3). Whilst the total station topographic survey reported by Graham
315 and Midgley (2000) took approximately 15 field-days, this aerial survey was
316 completed in 3 days and also provided high-resolution aerial imagery. The
317 standalone sUAS survey could have feasibly been completed in one day, however
318 this survey needed to be tied into the arbitrary coordinate system and datum used by
319 Graham and Midgley (2000). Operation of the sUAS is unfortunately restricted to dry
320 conditions, with relatively low wind speeds ($< 8 \text{ ms}^{-1}$). Despite specific weather
321 requirements, multi-rotor based systems appear to be well-suited to mountain
322 settings. They can be deployed where there is limited space for take-off and landing,
323 and offer a high-degree of control, which is beneficial when surveying in close
324 proximity to steep slopes. Regardless of the sUAS platform used for image
325 acquisition, the technique lends itself to surveying unstable or inaccessible terrain
326 where traditional survey methods would be unfeasible or unsafe.

327

328 UAS based image acquisition has clear benefits over existing full-scale airborne
329 image acquisition as the low survey altitude circumvents much of the weather
330 dependency (particularly cloud coverage) that affect full-scale airborne surveys
331 (Baltsavias, 1999). UAS also have the additional co-benefits of being less costly to
332 deploy in comparison to full scale airborne surveys and have the ability to produce
333 data products that are more scale appropriate for micro topographic investigations
334 than those provided by airborne LiDAR (Laliberte and Rango, 2009; Anderson and

335 Gaston, 2013). However, application of the SfM technique may be limited in some
336 geomorphological environments due to the presence of texturally 'smooth' or
337 reflective surfaces (e.g. snow cover or sand) which prohibit the extraction of
338 meaningful topographic data (Fonstad et al., 2013). Further work to investigate the
339 performance of automated image alignment over more texturally homogenous
340 surfaces may be beneficial where GPS information for camera positions are not
341 available. Care must be taken when acquiring coordinates for the GCPs used during
342 the image processing stage, due to the potential for erroneous readings to propagate
343 through the various derivative data products. Providing that the GCPs are accurately
344 surveyed, the automated nature of the approach is beneficial as it reduces the
345 potential for unintentional random error (e.g. as found to occur in the total station
346 dataset).

347

348 Although the production of a DSM from a dense point cloud produced a sub-
349 decimetre DSM, a coarser DSM (0.258 m per pixel) can be reconstructed from a
350 sparse point cloud of 2 million points with comparable error to that derived from a
351 dense point cloud of 30 million points. Where computational resources for both
352 image processing and data handling are limited or where data are not required at
353 sub decimetre resolution, producing DSMs from lower point densities maybe
354 desirable. The DSM presented here required ~ 7 h to point match and align the 543
355 images. An additional 43 min of processing time was needed to derive the dense
356 point cloud. Research to investigate the influence of point cloud density and the
357 resulting DSM error merits further investigation, although all DSMs should be
358 regarded as an abstraction, with some associated uncertainty (Fisher and Tate,
359 2006; Wechsler, 2007)

360 *5.3 sUAS–SfM as a tool for geomorphological mapping and monitoring*

361 *morphometric change*

362 The sUAS–SfM based approach appears to be a useful research tool that aids the
363 production of accurate geomorphological maps. A variety of data sources can be
364 used to compile geomorphological maps (Oguchi et al., 2011), with remotely sensed
365 data often requiring ground-truthing to ensure that landforms are accurately
366 recognised within a study area (Hubbard and Glasser, 2005; Knight et al., 2011).
367 From this perspective the recent availability of high-resolution airborne LiDAR
368 datasets are seen to be beneficial for the production of more accurate
369 geomorphological maps (Jones et al., 2007; Bishop et al., 2012), yet the limited
370 coverage of LiDAR surveys mean researchers do not always have access to high-
371 resolution data. In such cases the sUAS–SfM approach could be utilised by
372 researchers who wish to produce their own ultra-high-resolution DSMs and
373 orthophotos to aid field-mapping campaigns. Researchers should determine whether
374 the spatial coverage offered by sUAS is useful for their investigation. Here, a
375 localised area of 0.211 km² was surveyed over four separate flights. This is unlikely
376 to be sufficient for all geoscientific applications, however as sUAS technology
377 improves, greater survey coverage per flight may be permitted.

378

379 A further application of sUAS–SfM based surveys is morphometric change detection
380 due to how readily the technique can be deployed for use. Quantification of
381 geomorphological change through the comparison of multi-temporal DEMs is a well-
382 established practice applied to a range of geomorphological settings (coastal, glacial,
383 hillslope, fluvial, etc.; e.g. Pyle et al., 1997; Schiefer and Gilbert, 2007; Dewitte et al.,

384 2008; Marzoff and Poesen, 2009; Mitasova et al., 2009; Hugenholtz, 2010; Irvine-
385 Fynn et al., 2011; Carrivick et al., 2012). In some cases quantifying morphometric
386 change can be problematic where the rate of change is below or close to the
387 achievable accuracy of a given topographic survey technique (Williams, 2012). SfM
388 integrated with sUAS based image acquisition has recently been used for change
389 detection. For example, Whitehead et al. (2013) successfully completed repeat SfM
390 surveys to report on the thinning and motion of Fountain Glacier (Alaska) over a one
391 year period, with the first survey utilising a fixed wing UAS for image acquisition.
392 Lucieer et al. (2013) also used the sUAS–SfM approach, comparing multi-temporal,
393 multi-rotor derived aerial images to monitor landslide displacements at sub-
394 decimetre accuracies. The now widespread availability of aerial platforms and SfM
395 packages adds the range of mapping and survey techniques available to
396 geomorphologists. The technique is a logical choice due to the achievable survey
397 accuracies and potential to monitor geomorphological change at smaller spatial
398 scales remotely.

399

400 **6 Conclusions**

401 The integrated use of sUAS and SfM technologies for the acquisition of sub-
402 decimetre resolution DSMs has been investigated. The technique is shown to be
403 superior to conventional total station survey in terms of resolution, time required for
404 data acquisition, and has the additional benefit of providing ultra-high-resolution
405 orthorectified aerial imagery. DSM spatial resolutions of 0.088 m were achieved from
406 an approximate flight altitude of 117 m AGL whilst using a consumer-grade 18 MP
407 digital camera. Unintentional random error on the total station dataset, vegetation

408 and steep terrain are shown to promote vertical disagreement between the two
409 datasets. Where vegetation is sparse, a vertical difference of 0.200 m RMSE was
410 achieved. Overall, the technique is shown to provide exceptionally high-resolution
411 topographic datasets and aerial imagery. The repeatability of the technique where
412 surveys can be benchmarked or georeferenced using dGPS could offer not only
413 unprecedented spatial resolutions, but also high temporal resolution for monitoring
414 on-going geomorphological processes in a range of environments.

415

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662 **Fig. 1** Maps showing the study site location in relation to (a) North-west Europe, and
663 (b) North Wales. ©Crown Copyright/database right 2014. An Ordnance
664 Survey/EDINA supplied service. (c) A ground-level panoramic photograph of the
665 moraine-mound complex which is located on both the left and right of Llyn Idwal.

666

667 **Fig. 2** A schematic drawing of the S800 hexacopter.

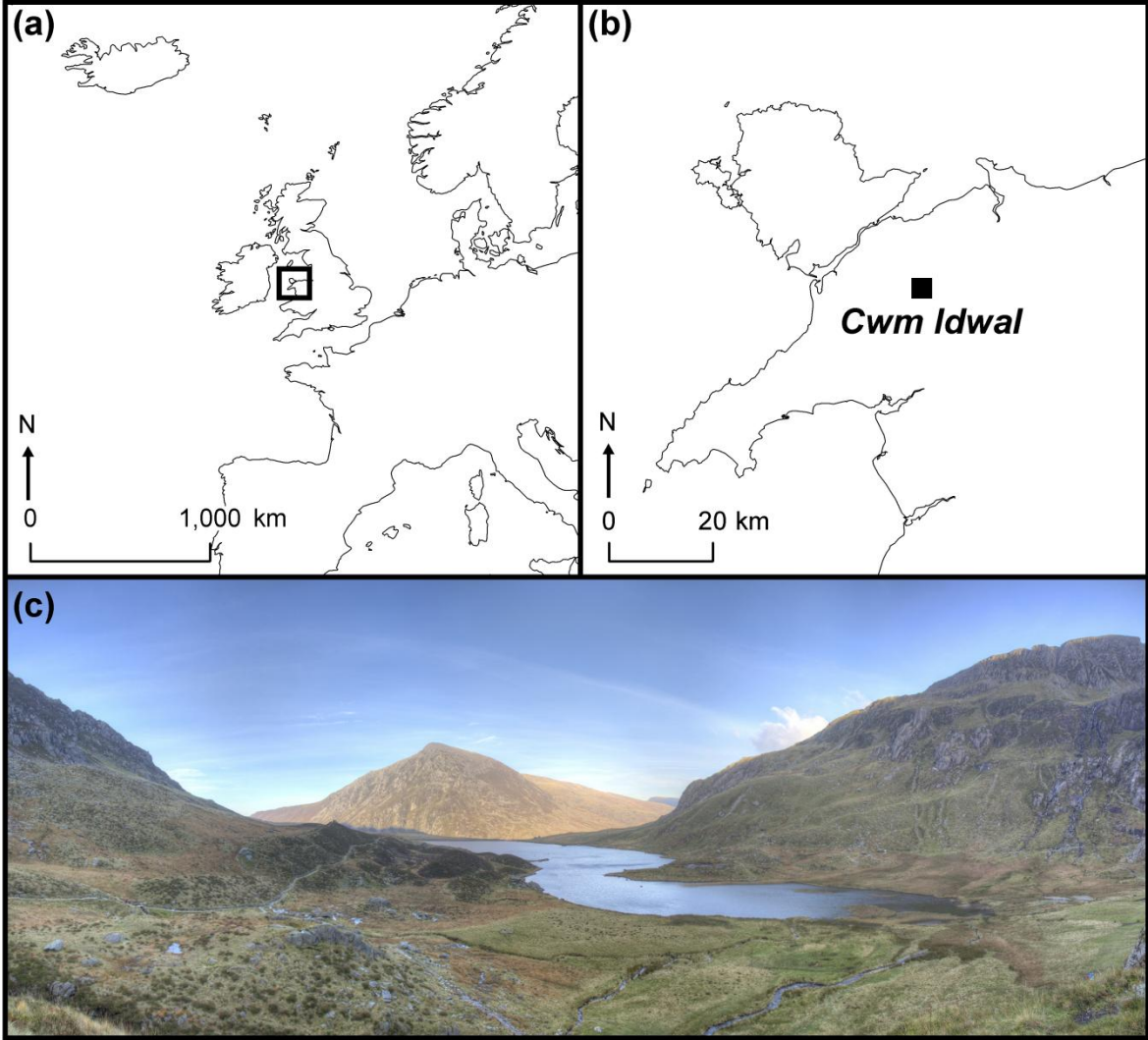
668

669 **Fig. 3** Maps displaying the topographic data and analysis of vertical disagreement.
670 (a) The distribution of 7761 ground-survey points and 19 SfM ground-control points
671 across the survey area. Two zones (Z1 and Z2) of distinct ground cover are
672 delimited. (b) A hillshaded DSM at 0.088 m per pixel resolution derived from the
673 sUAS–SfM survey. (c) A raster surface of vertical difference produced using an
674 ordinary kriging function at a resolution of 2.1 m per pixel. The spatial extent of the
675 spot heights is delimited by the dashed line.

676

677 **Fig. 5** The occurrence of vertical difference in association with: (a) vegetation, (b)
678 near vertical and in places partially overhanging bedrock rafts, and (c) positional
679 misregistration close to near vertical slopes

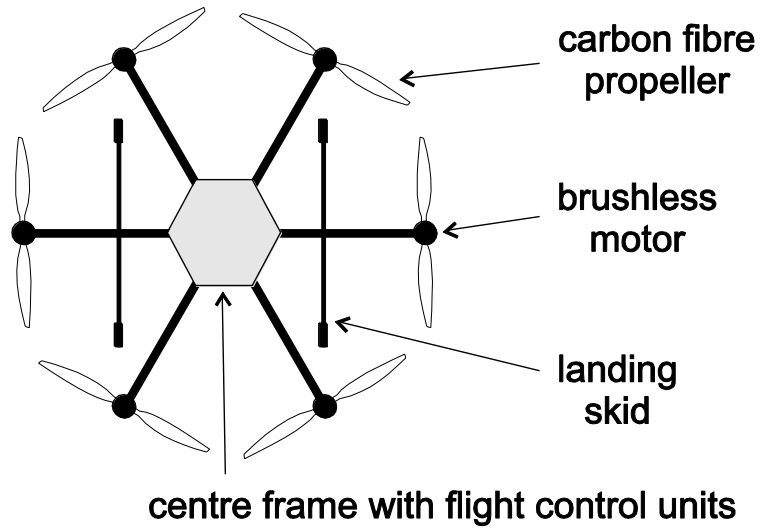
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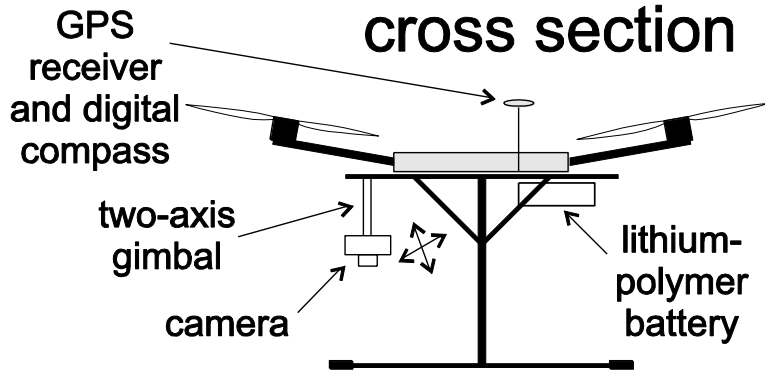
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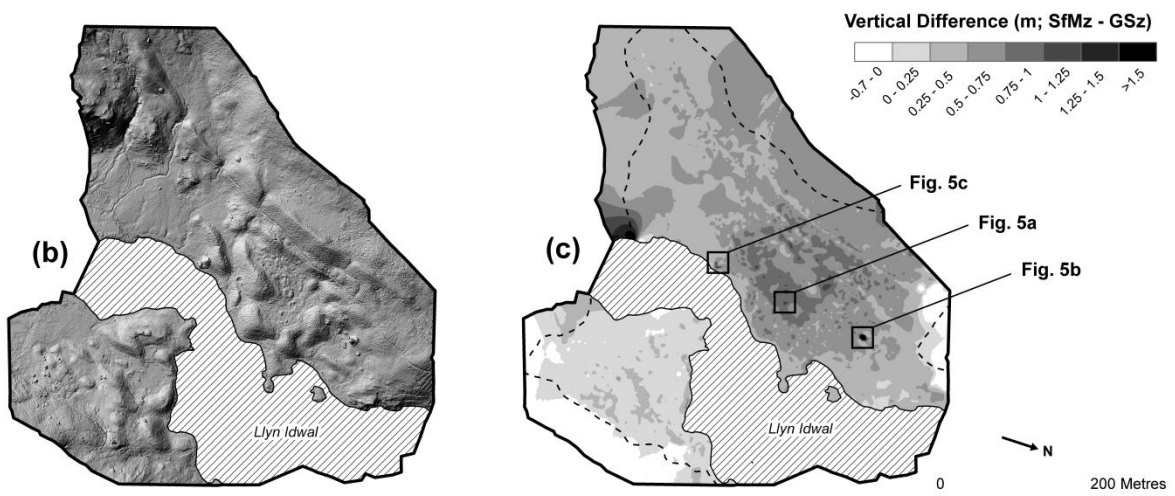
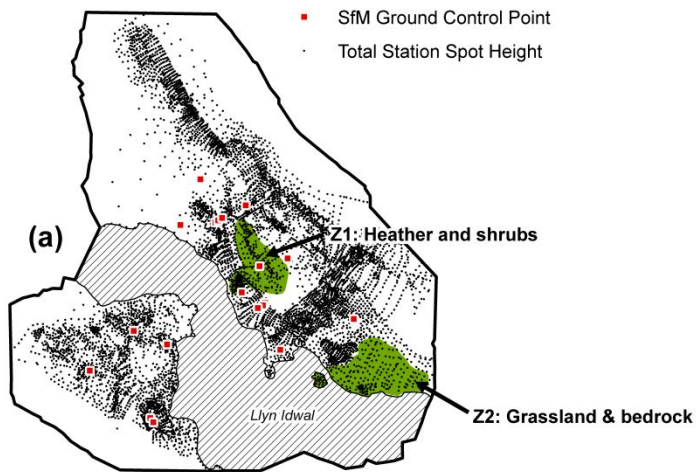
plan view

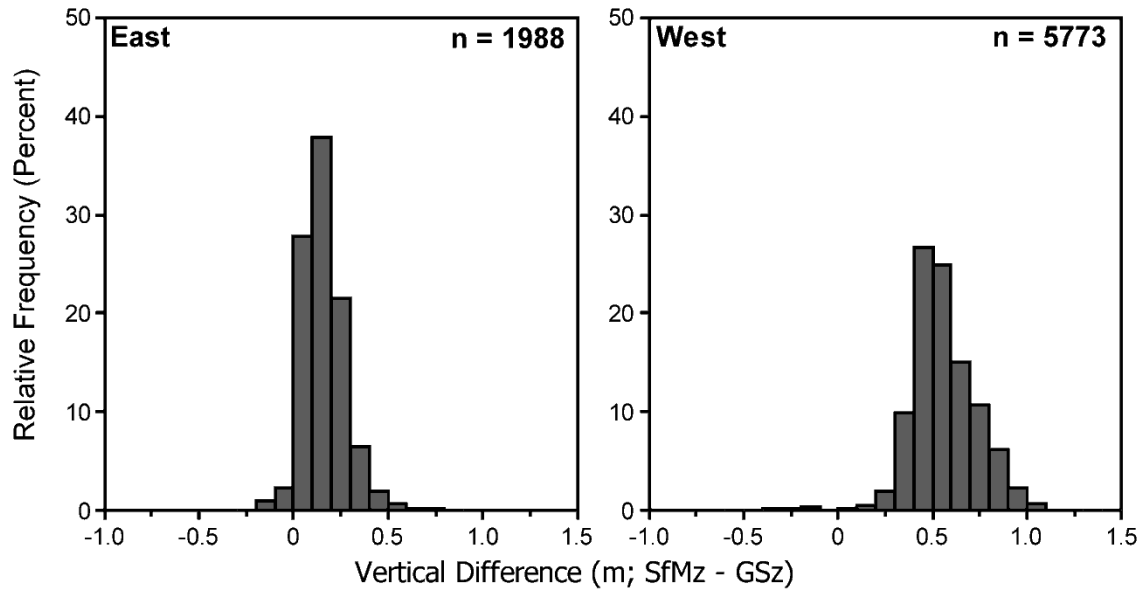


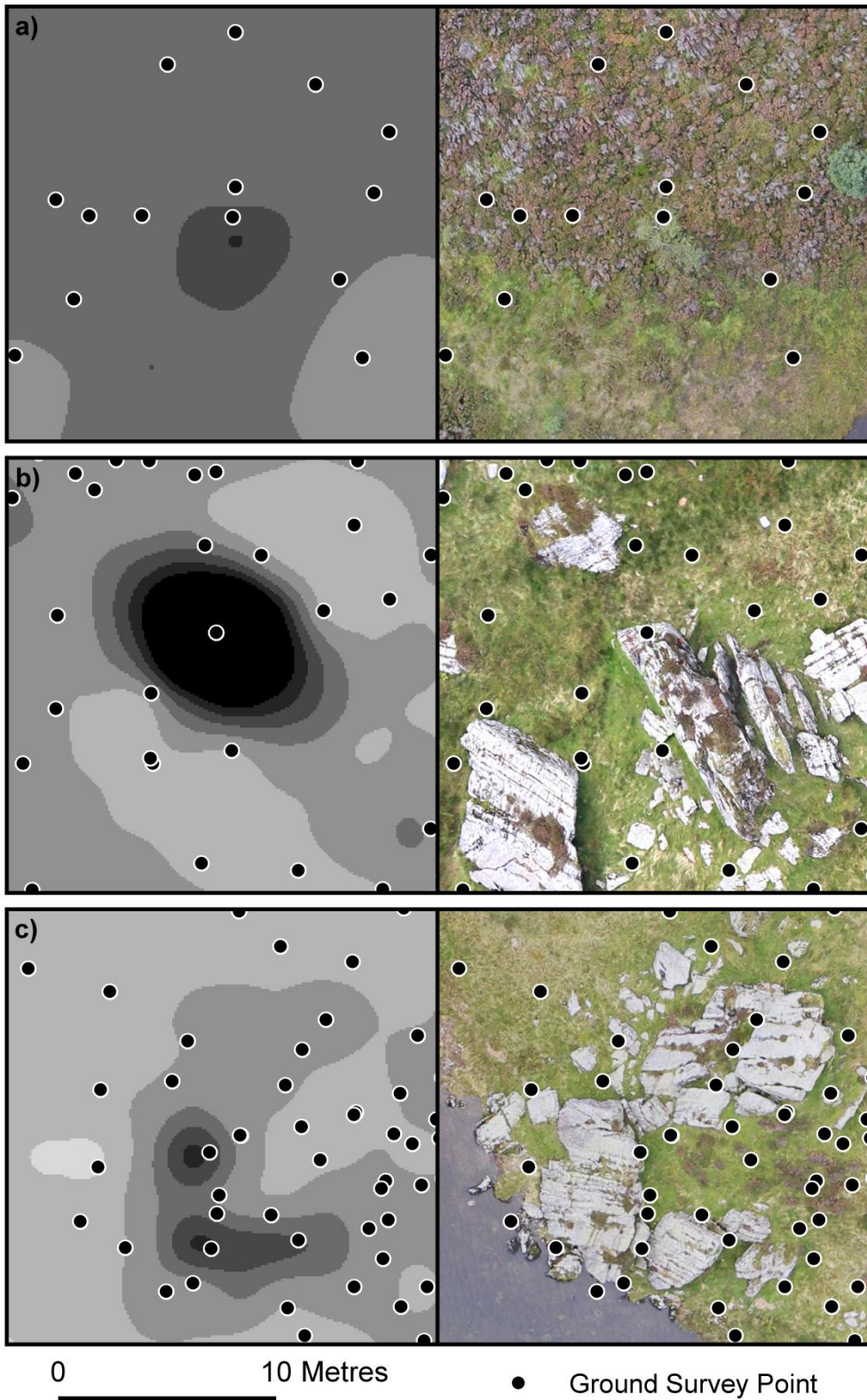
cross section



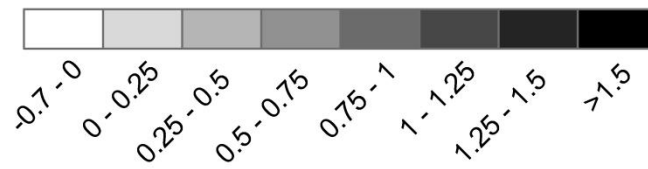
approximate scale guide 0 50cm







Vertical Difference (m; SfMz - GSz)



687 **Table 1** Statistics for the vertical difference between the Cwm Idwal topographic
 688 datasets

<i>Area</i>	<i>Total Observations (n)</i>	<i>RMSE</i>	<i>MD^a</i>	<i>RMSE (<20°)</i>	<i>RMSE (≥20°)</i>
All	7761	0.517	0.454	0.468 (n = 4527)	0.578 (n = 3234)
East	1988	0.200	0.155	0.169 (n = 1306)	0.247 (n = 682)
West	5773	0.588	0.557	0.544 (n = 3222)	0.639 (n = 2551)
Z1	244	0.796	0.820	0.789 (n = 102)	0.821 (n = 142)
Z2	205	0.362	0.341	0.354 (n = 152)	0.384 (n = 53)

689 ^a Mean Difference

690 **Table 2** Calculated RMSE for vertical difference binned by slope angle.

<i>Bin</i>	<i>RMSE</i>	<i>Observations (n)</i>
0 – 9	0.444	1864
10 – 19	0.482	2662
20 – 29	0.543	1967
30 – 39	0.603	952
40 – 49	0.678	263
50 – 59	0.739	36
60 – 69	0.729	10
70 – 79	0.838	6
80 – 90	2.222	1

691

692

693 **Table 3** Comparative table of known vertical differences between SfM topographic
 694 surveys and various validation datasets in a range of geomorphological
 695 environments

Study	Geomorphological setting	Camera	Camera Platform	Survey Altitude (m AGL)	Validation Data	Vertical Difference
Westoby et al. (2012)	Coastal	SLR: Model not specified	None	Ground-level	Terrestrial Laser Scanner	94% points values within +/- 1 m
Hugenholtz et al. (2013)	Aeolian	Olympus PEN Mini E-PM1	Fixed-wing sUAS	200	RTK GPS	RMSE = 0.29 m
Fonstad et al. (2013)	Exposed Bedrock/Fluvial	Canon A480	Helikite	10-70	LiDAR	RMSE = 1.05 m
Javernick et al. (2014)	Fluvial (Braided Channel)	Canon (10.1 MP): Model not specified.	Full-scale helicopter	600-800	RTK GPS	RMSE = 0.13 – 0.37 m
This study	Glacial landforms (Vegetated)	Canon EOS-M (18 MP)	Multi-rotor sUAS	117 (average)	Total Station	RMSE = East: 0.200 m West: 0.588 m All: 0.517 m

696